

垂直线源灌土壤湿润体尺寸预测模型研究

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摘要: 基于 HYDRUS-2D 模型建立了垂直线源灌土壤水分运动数学模型, 设置 81 种情景, 模拟获得不同土壤质地、初始含水率以及线源长度、线源直径和埋深条件下的湿润体变化过程。湿润体尺寸主要受土壤质地影响, 土壤质地越粗, 湿润锋运移越快, 线源长度、线源直径和埋深对其影响较小。土壤湿润锋运移过程符合幂函数关系, 幂函数指数在水平和垂直向上方向上变化较小, 而在垂直向下方向上随饱和导水率(K_s)的增大而增大; 幂函数系数随 K_s 的增大而增大。提出了包含 K_s 在内的垂直线源灌土壤湿润体尺寸预测模型, 试验验证了所建模型的可靠性, MAE 和 RMSE 接近 0, PBIAS 在 -4% ~ 9% 之间, NSE 不小于 0.929, 说明预测效果良好。所建模型仅需 K_s 即可推求, 试验设计简单, 初步实现了由土壤物理参数预测垂直线源灌土壤湿润锋运移距离的可能。

关键词: 垂直线源灌; 湿润体尺寸; 经验模型; HYDRUS-2D 模型

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Empirical Model for Predicting Wetted Soil Dimensions under Vertical Line Source Irrigation

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Abstract: The wetting pattern is difficult to observe during the vertical line source irrigation. Knowledge of the dimensions of wetted soil around the emitter under irrigation is essential to design of cost-effective and efficient vertical line source irrigation systems. Based on the HYDRUS-2D model, a mathematical model of soil water movement in vertical line source irrigation was established. And 81 scenarios were set up to simulate the changing process of wetted soil under different soil textures, initial water content, line source length, diameter and depth conditions. The dimensions of wetted soil were mainly affected by soil texture, the thicker the soil texture was, the faster the wetted front was moved, and the length, diameter and buried depth of line source had little influence on it. The migration process of soil wetting front was in accordance with the power function relationship. The power function index was changed little in the horizontal and vertical upward directions, but it was increased with the saturated hydraulic conductivity (K_s) in the vertical downward direction. The power function coefficient was increased with the increase of K_s . An empirical model for predicting the wetted soil dimensions under vertical line source irrigation containing K_s was proposed, the model reliability was verified by using experimental data. MAE and RMSE were close to 0, PBIAS was between -4% and 9%, and NSE was not less than 0.929, indicating that the prediction effect was good. The model can be estimated only by K_s , and the experimental design was simple. The possibility of predicting the soil wetting front migration distance under vertical line source irrigation by soil physical parameters was initially realized.

Key words: vertical line source irrigation; wetted soil dimension; empirical model; HYDRUS-2D model

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0 引言

我国西北旱区土地资源丰富,日照充足,昼夜温差大,是发展林果业的理想区域^[1]。然而,这些地区降雨量有限,果树生产在很大程度上取决于灌溉^[2]。传统灌溉方式耗水量大、水分利用效率低,不利于生态经济的可持续发展^[3-4]。因此,灌溉技术和水资源管理的改进将发挥重要作用。

垂直线源灌是一种适用于深根植物的节水灌溉方法,其灌水器垂直埋入土体,灌水过程中,水分直接进入植物根部,湿润体不易观测^[5]。湿润体的形状及大小影响着植物的生长与产量,了解垂直线源灌湿润体动态变化特征,可确保灌水器在活性根区的精确放置,对设计经济高效的垂直线源灌溉系统至关重要^[6-7]。土壤质地是决定灌溉设计参数的重要因素,地下灌溉系统的设计应考虑土壤地质的影响^[8-12];相同土壤质地条件下,土壤容重增加,孔隙度减小,土体入渗能力下降^[13-16];土壤初始含水率决定了渗透初期的土壤水势,土壤初始含水率增加,湿润体尺寸逐渐增大^[5,11,17-20]。线源长度和直径决定了灌水器渗水界面的大小,线源长度或直径增大,渗水界面面积变大,意味着水分进入土壤的通道增加,导致相同时间内入渗水量增多,湿润体随之增大^[21-24]。因此,从实用角度出发,线源直径既要尽可能的大以加快其渗水速率,提高灌水均匀度,又要尽可能的小以减弱对作物根系生长的影响。灌水器埋深直接改变湿润体水分分布位置,是实现作物根系与湿润体有效匹配的关键因素,埋深过浅会增加地表水分无效蒸发,埋深过深又会引起深层渗漏和表土水分亏缺^[25-29]。因此,埋深应与土壤条件、根系分布及耕作要求等相适应。灌溉必须适时适量,灌水时间过早或过晚、灌水定额过大或过小都是无益的。DU 等^[2]研究表明,中国西北干旱区苹果根区土壤含水率低于田间持水率的 50%~55% 时,会对树木生长和最终产量造成水分胁迫;周罕觅等^[30]研究水肥耦合对 3 年生苹果幼树生长、产量、品质及水肥利用的效应,得出灌水下限为田间持水率的 65%~75%;贾俊杰等^[31]指出,SH 矮砧苹果幼树滴灌条件下适宜灌水下限为 60% 的田间持水率。孙三民等^[32]通过小区试验,确定 13 L/(棵·次)的灌水量为适宜的新疆红枣间接地下滴灌灌溉模式;张陆军等^[33]指出,陕北山地梨枣树涌泉根灌时,每株 2 个灌水器及每个灌水器 40 L/(株·次)的灌水量组合是适宜的布置方式;吴悠等^[34]通过遮雨棚下可称量式蒸渗桶试验得出,生育期内 8.4 L/(株·次)为柱状苹果树相对节水的灌溉模式。

湿润体动态变化可通过湿润锋距离量化表征。国内外学者开发了一些用于确定湿润锋距离的模型,其中最常见的是分析模型^[21,35-38]、数值模型^[39-40]和经验模型^[11,41-45]。通常,通过求解特定初始和边界条件的控制方程(Richards 方程)来开发分析模型和数值模型,而使用实验或数值模拟的回归分析来开发经验模型。文献[46-48]对数值和经验模型进行了比较和评估,研究表明,HYDRUS 模型计算结果能较好地反映土壤水分运动基本规律,但模型较复杂,需输入大量参数才能模拟计算;另外,每个经验模型都是土壤水力特性和灌溉参数的函数方程,形式较简单,但仅适用于具体的灌溉技术,如开发的滴灌或沟灌湿润体预测模型并不适用于垂直线源灌。因此,有必要开发一种可以预测垂直线源灌土壤湿润体尺寸的经验模型,为确定适宜的灌水技术参数和实现灌溉系统优化运行提供实用而方便的手段。

数值模拟方法可对不同土壤特性、不同灌水器规格和不同设计参数条件下的土壤水分运动过程进行模拟^[49-51]。李淑芹等^[52]、FAN 等^[53]通过试验验证了垂直线源灌 HYDRUS-2D 模拟结果的有效性。基于此,本文采用 HYDRUS-2D 软件,模拟研究土壤质地、初始含水率、线源长度、线源直径和线源埋深对垂直线源灌湿润体运移特征值的影响;利用模拟数据筛选影响湿润体运移的主导因素,进而构建预测湿润体尺寸的简化经验模型;最后,通过土箱试验验证经验模型的可靠性。

1 材料与方法

1.1 试验设计

试验装置由 3 部分组成:土箱、马氏瓶和线源灌水器,如图 1 所示。土箱由 10 mm 厚有机玻璃制成,长×宽×高为 50 cm×50 cm×100 cm。土箱底部留有多多个通气孔(直径 2 mm),以防气阻发生。线源与土箱接触面开取土孔(直径 2 cm,间距 5 cm),用于测量灌溉结束时的土壤含水率。线源采用 1/4 圆柱体,底端密封,管底向上 l 长度的柱面均匀开孔。马氏瓶直径为 10 cm,高度为 100 cm。试验前,将供试土样按设定的初始含水率加水,均匀混合后,用塑料薄膜密闭静置 1 d,待土壤水分分布均匀后,按设计容重分层(5 cm)装入土箱,以获得均匀土壤剖面。为了便于观察土壤湿润体变化过程,将线源灌水器用纱布包裹,并置于土箱一角,确保灌水器管壁与土壤紧实接触,待次日进行入渗试验。试验中,马氏瓶提供恒定水头,按先密后疏的时间间隔记录累积入渗量,并用马克笔绘制湿润锋运移图。入渗达到设

定灌水定额后停止供水,迅速从灌水器两侧预留孔取土,用干燥法(105℃干燥 24 h)测定土壤含水率。为尽量消除试验误差,每个试验重复 3 次。

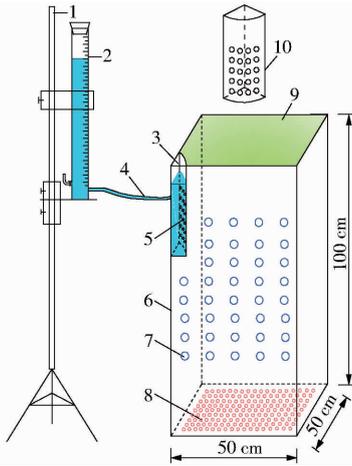


图 1 垂直线源灌试验装置及灌水器细部结构

Fig. 1 Experimental equipment for vertical line source irrigation and detailed structure of emitter

1. 调节水头支架 2. 马氏瓶 3. 灌水器 4. 橡胶管 5. 渗水孔
6. 土箱 7. 取土孔 8. 通气孔 9. 土壤表面 10. 灌水器细部结构

参照文献[2, 30-34]研究成果,取民勤地区砂壤土(容重 $\gamma_d = 1.45 \text{ g/cm}^3$, 田间持水率 $\theta_f = 0.332 \text{ cm}^3/\text{cm}^3$, 饱和导水率 $K_s = 0.039 \text{ cm/min}$)和风沙土($\gamma_d = 1.56 \text{ g/cm}^3$, $\theta_f = 0.051 \text{ cm}^3/\text{cm}^3$, $K_s = 0.345 \text{ cm/min}$),每种土壤采用 2 种处理(初始含水率 $\theta_0 = 60\% \theta_f$ 、线源直径 $d = 4 \text{ cm}$ 、线源长度 $l = 20 \text{ cm}$ 、线源埋深 $b = 40 \text{ cm}$ 、灌水量 $V = 40 \text{ L}$; $\theta_0 = 70\% \theta_f$ 、 $d = 6 \text{ cm}$ 、 $l = 30 \text{ cm}$ 、 $b = 50 \text{ cm}$ 、 $V = 40 \text{ L}$)进行垂直线源灌土壤入渗试验。

1.2 数学建模

1.2.1 基本方程

假设土壤是均匀和各向同性的,垂直线源灌可概念化为轴对称的三维入渗过程。使用 HYDRUS-2D 模拟^[54]。土壤水分运动控制方程为 Richards 方程

$$\frac{\partial \theta}{\partial t} = \frac{1}{r} \frac{\partial}{\partial r} \left(rK(h) \frac{\partial h}{\partial r} \right) + \frac{\partial}{\partial z} \left(K(h) \frac{\partial h}{\partial z} \right) - \frac{\partial K(h)}{\partial z} \quad (1)$$

式中 r ——径向坐标值, cm

z ——垂向坐标值, cm, 向下为正

θ ——土壤含水率, cm^3/cm^3

h ——压力水头, cm

t ——时间, min

$K(h)$ ——土壤非饱和导水率, cm/min

采用 van Genuchten - Mualem (VG-M) 方程^[55-56]描述土壤水分特征曲线和非饱和导水率

$$S_e = \frac{\theta - \theta_r}{\theta_s - \theta_r} = \frac{1}{(1 + |\alpha h|^n)^m} \quad (2)$$

$$K(h) = K_s S_e^{0.5} [1 - (1 - S_e^{1/m})^2] \quad (3)$$

其中

$$m = 1 - 1/n$$

式中 S_e ——土壤相对饱和度

θ_r ——土壤残余含水率, cm^3/cm^3

θ_s ——土壤饱和含水率, cm^3/cm^3

α ——与进气值成反比的经验参数, cm^{-1}

n, m ——影响土壤水分特征曲线形状的经验常数

1.2.2 定解条件

图 2(图中 A、B、C 分别为线源最高点、最低点和中心点)为本研究中用于模拟不同建模情景的初始和边界条件。

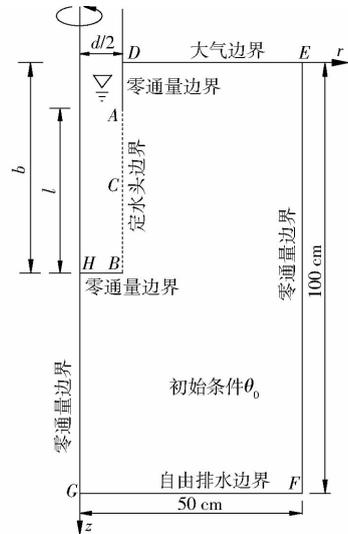


图 2 具有初始和边界条件的计算域

Fig. 2 Computational domain with initial and boundary conditions

在所有的模拟情景中,土壤含水率按初始含水率设置;上边界 DE 受大气条件影响,考虑灌水过程中地表为干土层,蒸发量很小,为简化计算,按零通量面设置;下边界 FG 不受灌水影响,为自由排水,按零通量面设置;左边界 GH 为灌水器中心轴,AD 为塑料管壁,均无水量交换,按零通量面设置;右边界 EF 灌溉水未到达,按零通量面设置;线源底部 BH 密封,为零通量边界;渗水面边界为充分供水方式,供水开始后很快达到饱和,可按定水头边界处理^[20,22]。

综上,初始条件可表述为

$$\begin{cases} \theta = \theta_0 \\ t = 0 \\ (r, z) \in \Omega \end{cases} \quad (4)$$

式中 θ_0 ——土壤初始含水率, cm^3/cm^3

Ω ——计算域(图 2)

边界条件可表述为

$$\begin{cases} -K(h) \left(\frac{\partial h}{\partial z} - 1 \right) = 0 & (t > 0, DE、FG \text{ 和 } BH \text{ 边界}) \\ \theta = \theta_s & (t > 0, AB \text{ 边界}) \\ -K(h) \frac{\partial h}{\partial r} = 0 & (t > 0, EF、GH \text{ 和 } DA \text{ 边界}) \end{cases} \quad (5)$$

表 1 HYDRUS 模拟中 9 种典型土壤的 VG-M 模型参数

Tab. 1 VG-M model parameters of nine typical soils in HYDRUS simulation

土壤质地	$\theta_r / (\text{cm}^3 \cdot \text{cm}^{-3})$	$\theta_s / (\text{cm}^3 \cdot \text{cm}^{-3})$	α / cm^{-1}	n	$K_s / (\text{cm} \cdot \text{min}^{-1})$
黏壤土*	0.095	0.410	0.019	1.31	0.0043
粉土*	0.034	0.460	0.016	1.37	0.0042
粉壤土*	0.067	0.450	0.020	1.41	0.0075
砂黏壤土*	0.100	0.390	0.059	1.48	0.0218
壤土*	0.078	0.430	0.036	1.56	0.0173
砂黏壤土**	0.029	0.430	0.079	1.87	0.0730
砂壤土*	0.065	0.410	0.075	1.89	0.0737
壤砂土*	0.057	0.410	0.124	2.28	0.2432
砂土*	0.045	0.430	0.145	2.68	0.4950

注: * 表示取自文献[57], ** 表示取自文献[52]。

不同质地土壤田间持水率采用 RAB 等^[58]建立的预测模型获得,具体表达式为

$$\theta_f = 8.05 + 1.68\theta_p - 1.62\theta_p^2 \quad (6)$$

式中 θ_p ——凋萎系数,可采用 VG-M 模型参数中的 θ_r 表示^[11], $\text{cm}^3 / \text{cm}^3$

1.2.4 求解方法

利用 HYDRUS-2D 进行数值求解。求解过程中,采用隐式差分格式进行时间离散,Galerkin 有限元法对土壤剖面进行空间离散。考虑到田间实际和计算精度的要求,确定有限单元计算域深度为 100 cm,宽度为 50 cm,空间步长为 1 cm,时间步长为 0.1 min,模拟历时由灌水定额(40 L)决定。

1.3 分析方法

垂直线源灌湿润体形状近似为“梨”型^[22,52]。选取 5 个特征值(A 点水平方向、B 点水平方向、C 点水平方向、C 点垂直向上和 C 点垂直向下)勾画出湿润体轮廓,点 A、B 和 C 见图 2。研究表明地下三维入渗土壤湿润锋运移过程可采用幂函数描述,且具有很高的精度^[22,59-61]。因此,采用幂函数定量分析垂直线源灌土壤湿润锋运移过程,其具体表达式为

$$U_C = l/2 + b_1 t^{a_1} \quad (7)$$

$$D_C = l/2 + b_2 t^{a_2} \quad (8)$$

$$R_A = d/2 + b_3 t^{a_3} \quad (9)$$

$$R_B = d/2 + b_4 t^{a_4} \quad (10)$$

$$R_C = d/2 + b_5 t^{a_5} \quad (11)$$

1.2.3 模拟方案

采用单因素分析法,设置 81 种情景,模拟分析不同土壤质地(表 1)、 θ_0 (50% θ_f 、60% θ_f 、70% θ_f)、 d (2、4、6 cm)、 l (10、20、30 cm) 和 b (30、40、50 cm) 等因素对垂直线源灌湿润体的影响。土壤质地 VG-M 模型参数取自 CARSEL 等^[57] 资料以及文献[52],如表 1 所示。

式中 $U_C、D_C$ ——C 点垂直向上、垂直向下湿润高度,cm

$R_A、R_B、R_C$ ——A、B、C 点湿润半径,cm

$a_1、a_2、a_3、a_4、a_5、b_1、b_2、b_3、b_4、b_5$ ——拟合参数

根据模拟结果分析湿润体运移规律,探讨影响机理,筛选主导因素,采用式(7)~(11)拟合获得 $a_1、a_2、a_3、a_4、a_5、b_1、b_2、b_3、b_4$ 和 b_5 值,研究拟合参数与主导因素间的量化关系,进而建立垂直线源灌土壤湿润体运移距离模型。

1.4 误差分析

选取 4 个指标,即平均绝对误差 MAE、均方根误差 RMSE、偏差百分比 PBIAS 和纳什效率系数 NSE,对湿润体尺寸的实测值和预测值进行误差分析。指标参数定义为

$$E_{\text{MAE}} = \frac{1}{N} \sum_{i=1}^N |M_i - S_i| \quad (12)$$

$$E_{\text{RMSE}} = \left[\frac{1}{N} \sum_{i=1}^N (M_i - S_i)^2 \right]^{0.5} \quad (13)$$

$$E_{\text{PBIAS}} = \frac{\sum_{i=1}^N (M_i - S_i)}{\sum_{i=1}^N M_i} \quad (14)$$

$$E_{\text{NSE}} = 1 - \frac{\sum_{i=1}^N (M_i - S_i)^2}{\sum_{i=1}^N (M_i - M_{\text{mean}})^2} \quad (15)$$

式中 E_{MAE} ——平均绝对误差

E_{RMSE} ——均方根误差
 E_{PBIAS} ——偏差百分比
 E_{NSE} ——纳什效率系数
 M_i ——第 i 个实测值
 S_i ——第 i 个模拟值
 M_{mean} ——实测值的平均值
 N ——数据总个数

MAE 和 RMSE 的数值越接近 0, PBIAS 为 $[-10, 10]$,

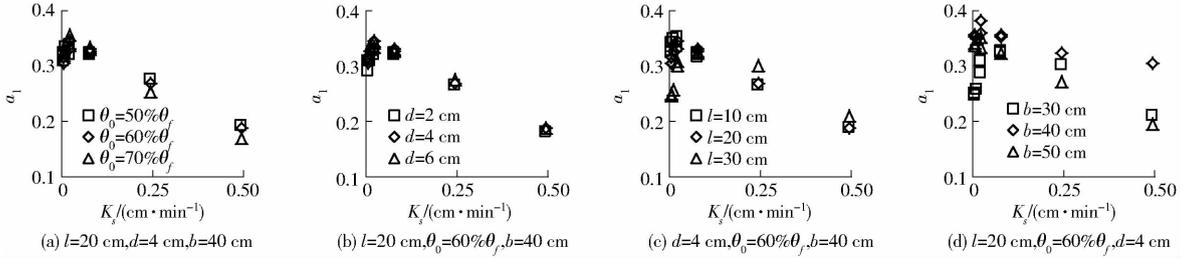


图3 拟合参数 a_1 随饱和导水率 K_s 的变化规律

Fig. 3 Variations of fitting parameter a_1 with saturated hydraulic conductivity K_s

由图3可见,相同 K_s 时,拟合参数 a_1 主要受线源长度和埋深的影响,而土壤初始含水率和线源直径对其影响较小。主要是线源长度和埋深影响湿润锋到达地表的时间,线源越长或埋深越浅,湿润锋到达地表时间越短,导致拟合参数 a_1 时数据量减少,产生了拟合误差。不同 K_s 情况下,拟合参数 a_1 随 K_s

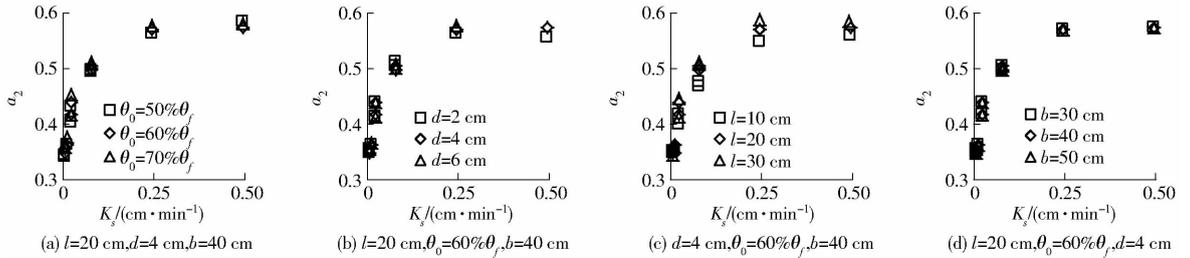


图4 拟合参数 a_2 随饱和导水率 K_s 的变化规律

Fig. 4 Variations of fitting parameter a_2 with saturated hydraulic conductivity K_s

长度和埋深对拟合参数 a_2 影响较小。拟合参数 a_2 随 K_s 的增大而增大,两者具有较好的幂函数关系,即 $a_2 = 0.65K_s^{0.108}$ ($R^2 = 0.996$)。

2.1.3 拟合参数 a_3

利用 HYDRUS-2D 模拟结果,采用式(9),拟合获得不同影响因素下 a_3 ,如图5所示。

由图5可见,土壤初始含水率、线源直径、线源

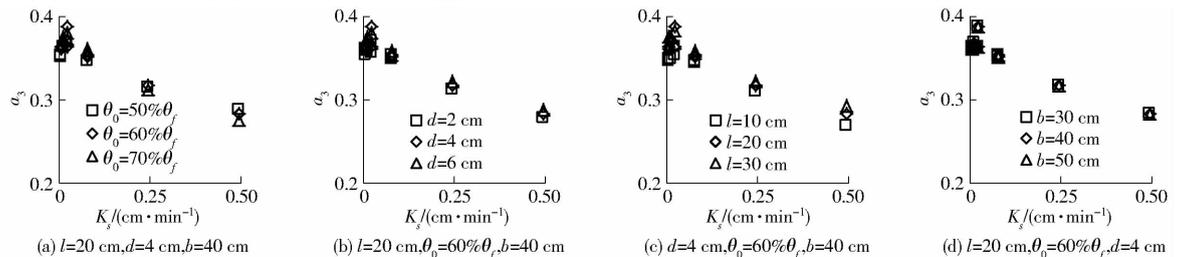


图5 拟合参数 a_3 随饱和导水率 K_s 的变化规律

Fig. 5 Variations of fitting parameter a_3 with saturated hydraulic conductivity K_s

NSE 越靠近 1, 表示模拟值与实测值差异越小,两者吻合越好^[62]。

2 结果与分析

2.1 拟合参数 $a_1 \sim a_5$ 影响因素分析

2.1.1 拟合参数 a_1

利用 HYDRUS-2D 模拟结果,采用式(7),拟合获得不同影响因素下 a_1 ,如图3所示。

先增大后减小,但增减幅度不大(0.172 ~ 0.355),为简化计算,可取平均值,即 $a_1 = 0.3$ 。

2.1.2 拟合参数 a_2

利用 HYDRUS-2D 模拟结果,采用式(8),拟合获得不同影响因素下 a_2 ,如图4所示。

由图4可见,土壤初始含水率和线源直径、线源

长度和埋深对拟合参数 a_3 影响较小。不同 K_s 情况下,拟合参数 a_3 随 K_s 先增大后减小,但增减幅度不大(0.269 ~ 0.381),为简化计算,可取平均值,即 $a_3 = 0.348$ 。

2.1.4 拟合参数 a_4

利用 HYDRUS-2D 模拟结果,采用式(10),拟合获得不同影响因素下 a_4 ,如图6所示。

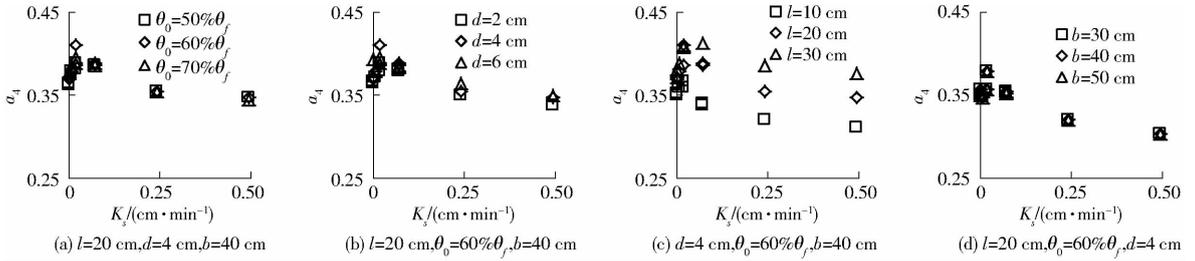


图 6 拟合参数 a_4 随饱和导水率 K_s 的变化规律

Fig. 6 Variations of fitting parameter a_4 with saturated hydraulic conductivity K_s

由图 6 可见, 相同 K_s 时, 拟合参数 a_4 主要受线源长度的影响, 而土壤初始含水率、线源直径和埋深对其影响较小。线源越长, 其渗水速率越快, 导致湿润体水分叠加效应增强, 同时, 由于重力势的作用, 下部湿润体的叠加效应强于上部。不同 K_s 情况下, 拟合参数 a_4 随 K_s 先增大后减小, 但增减幅度不大

(0.298 ~ 0.414), 为简化计算, 可取平均值, 即 $a_4 = 0.374$ 。

2.1.5 拟合参数 a_5

利用 HYDRUS-2D 模拟结果, 采用式 (11), 拟合获得不同影响因素下 a_5 , 如图 7 所示。

由图 7 可见, 类似于拟合参数 a_4 , 相同 K_s 条件

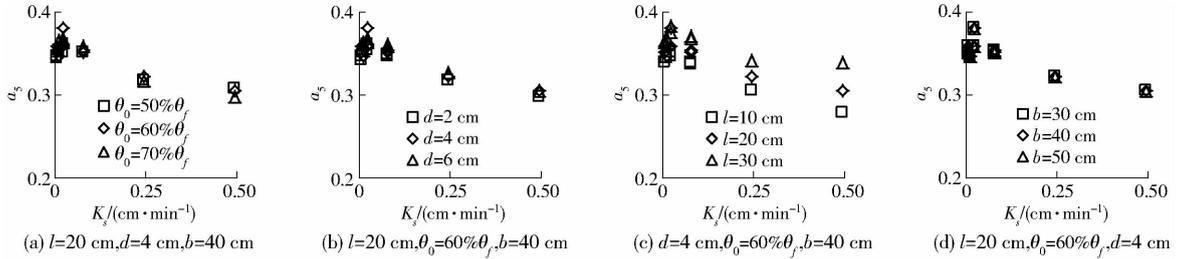


图 7 拟合参数 a_5 随饱和导水率 K_s 的变化规律

Fig. 7 Variations of fitting parameter a_5 with saturated hydraulic conductivity K_s

下, 拟合参数 a_5 主要受线源长度的影响, 不同 K_s 情况下, 拟合参数 a_5 随 K_s 先增大后减小, 但增减幅度不大 (0.278 ~ 0.379), 为简化计算, 可取平均值, 即 $a_5 = 0.346$ 。

源灌湿润体尺寸简化预测模型, 即

$$U_c = l/2 + 3.3198K_s^{0.157}t^{0.300} \quad (21)$$

$$D_c = l/2 + (5.0952K_s + 1.0482)t^{0.65K_s^{0.108}} \quad (22)$$

$$R_A = d/2 + 4.036K_s^{0.2174}t^{0.348} \quad (23)$$

$$R_B = d/2 + 5.0938K_s^{0.3004}t^{0.374} \quad (24)$$

$$R_C = d/2 + 5.1328K_s^{0.2491}t^{0.346} \quad (25)$$

综上所述, 拟合参数 a_1 、 a_3 、 a_4 和 a_5 随 K_s 、 θ_0 、 d 、 l 和 b 的变化规律不明显, 且变化幅度较小。为简化计算, 分别取其平均值。 a_2 随 θ_0 、 d 、 l 和 b 的增减而稍有变化, 但变化较小, 而随 K_s 的增大而增大, 两者具有较好的幂函数关系。

2.3 模型验证

利用试验数据对简化预测模型进行验证。将砂壤土和风沙土的 K_s 分别代入式 (21) ~ (25), 得 2 种土壤的垂直线源灌湿润体尺寸简化预测模型为:

$$\text{民勤砂壤土} \begin{cases} U_c = l/2 + 1.99t^{0.300} \\ D_c = l/2 + 1.25t^{0.458} \\ R_A = d/2 + 2.00t^{0.348} \\ R_B = d/2 + 1.92t^{0.374} \\ R_C = d/2 + 2.29t^{0.346} \end{cases} \quad (26)$$

$$\text{民勤风沙土} \begin{cases} U_c = l/2 + 2.90t^{0.300} \\ D_c = l/2 + 3.20t^{0.592} \\ R_A = d/2 + 3.35t^{0.348} \\ R_B = d/2 + 3.93t^{0.374} \\ R_C = d/2 + 4.14t^{0.346} \end{cases} \quad (27)$$

将 $a_1 \sim a_5$ 分别代入式 (7) ~ (11), 得

$$U_c = l/2 + b_1t^{0.300} \quad (16)$$

$$D_c = l/2 + b_2t^{0.65K_s^{0.108}} \quad (17)$$

$$R_A = d/2 + b_3t^{0.348} \quad (18)$$

$$R_B = d/2 + b_4t^{0.374} \quad (19)$$

$$R_C = d/2 + b_5t^{0.346} \quad (20)$$

2.2 拟合参数 $b_1 \sim b_5$ 影响因素分析

利用 HYDRUS-2D 模拟结果, 采用式 (16) ~ (20), 再次拟合获得不同 θ_0 、 d 、 l 和 b 组合下的 $b_1 \sim b_5$, 如图 8 所示。

由图 8 可见, 拟合参数 b_1 、 b_2 、 b_3 、 b_4 和 b_5 均随 K_s 的增大而增大, θ_0 、 d 、 l 和 b 对其影响相对较小。进一步分析发现 b_1 、 b_3 、 b_4 和 b_5 与 K_s 具有很好的幂函数关系, 而 b_2 与 K_s 呈线性关系。基于此, 可得垂直线

将简化预测模型计算值与试验特征值 (2 个处理, 3 个重复) 进行对比分析, 如图 9 所示。

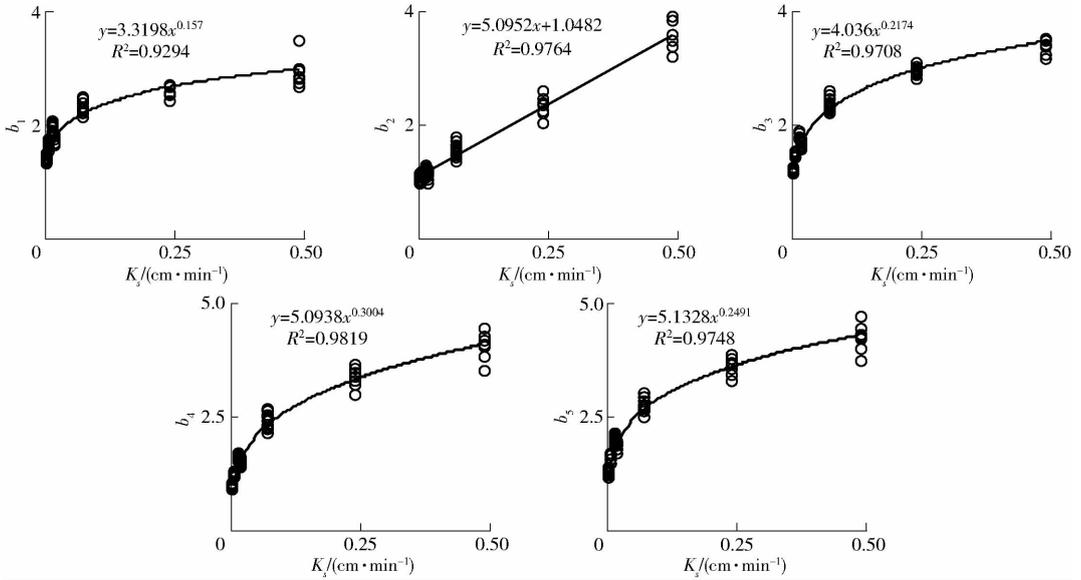


图8 拟合参数 $b_1 \sim b_5$ 随饱和导水率 K_s 的变化规律

Fig. 8 Variations of fitting parameter $b_1 \sim b_5$ with saturated hydraulic conductivity K_s

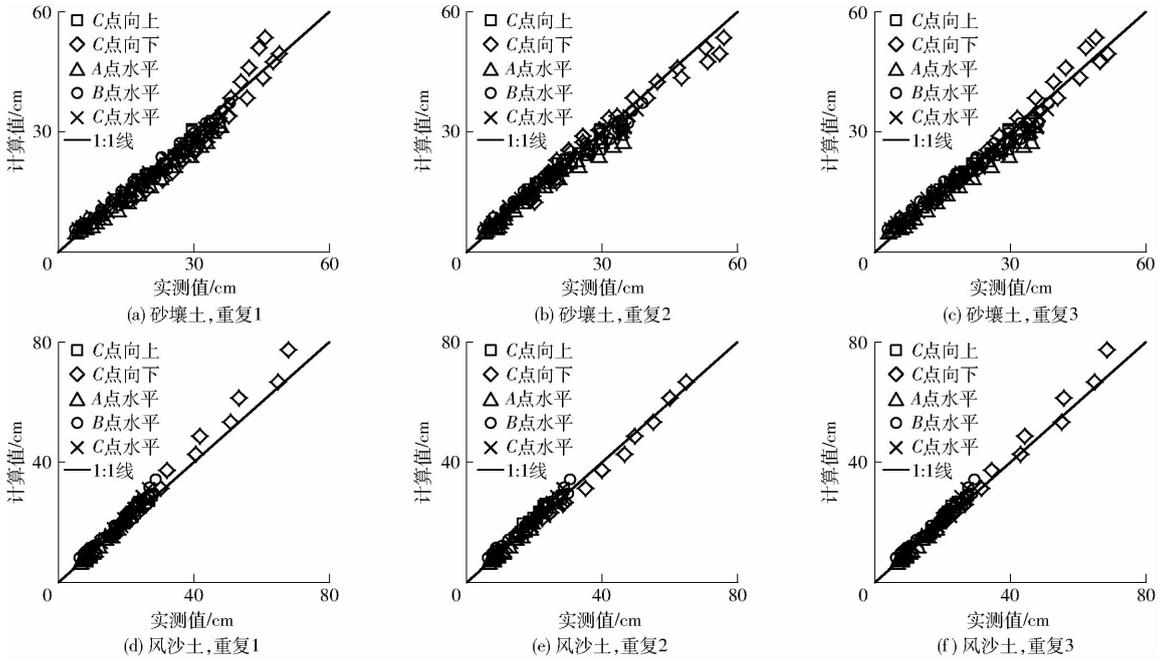


图9 2种土壤垂直线源灌土壤湿润体尺寸计算值与实测值对比

Fig. 9 Comparisons of calculated and measured values of wetted soil dimensions for two soil types under vertical line source irrigation

采用式(12)~(15),对计算值与实测值进行统计分析,计算结果见表2。

由表2可知,MAE和RMSE接近0,PBIAS为-4%~9%之间,NSE靠近1(NSE不小于0.929),说明简化预测模型计算值与实测值一致性良好,但仍存在一定误差,究其原因可能是垂直线源灌土壤湿润锋运移距离受土壤质地影响最大,而土壤初始含水率以及线源直径、长度和埋深对其尚有一些影响,为了简化计算,仅考虑了土壤质地的影响,建立了单变量模型,从而在一定程度上影响了计算结果的准确性。另外,仅采用饱和导水率 K_s 来表征不同

土壤质地湿润锋运移规律也是存在部分误差的原因之一。

表2 计算值与实测值统计分析

Tab. 2 Statistical analysis of calculated and measured values

土壤质地	编号	MAE/cm	RMSE/cm	PBIAS/%	NSE
民勤砂壤土	重复1	2.337	2.875	8.471	0.929
	重复2	1.663	2.348	4.353	0.952
	重复3	1.900	2.468	5.133	0.947
民勤风沙土	重复1	1.161	1.970	-3.996	0.977
	重复2	0.956	1.364	0.704	0.987
	重复3	0.978	1.602	-2.397	0.985

3 结论

(1)垂直线源灌湿润体尺寸主要受土壤质地影响,土壤质地越粗(K_s 越大),湿润锋运移越快;线源长度、线源直径和埋深对其影响较小。

(2)土壤湿润锋运移过程符合幂函数关系,幂函数系数随 K_s 的增大而增大,幂函数指数在垂直向

上和水平方向上变化较小,而在垂直向下方向上随 K_s 的增大而增大。

(3)提出了包含 K_s 的垂直线源灌土壤湿润体尺寸预测模型,利用试验验证了预测模型的有效性,初步实现了由土壤物理参数预测垂直线源灌土壤湿润锋运移距离的可能。

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